

Remotely sensed resilience of tropical forests

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13 1 Data sets and sampling strategy

14 1.1 Optical and microwave satellite data

15 Optical MODIS Normalized Difference Vegetation Index (NDVI) is derived from the bidirec-
16 tional reflectance distribution function (BRDF) adjusted reflectance measurements (MCD43C4
17 version 5, 2000–2011, 5.6km). The reflectance product MCD43C4 provides reflectance data ad-
18 justed using a BRDF to model the values as if they were taken from nadir view to account for
19 potential artefacts introduced by varying sun-sensor geometry¹. Both Terra and Aqua MODIS
20 sensor data are used in the generation of this product, providing the highest probability for qual-
21 ity input data and designating it as an MCD, meaning Combined, product. Fill values and mea-
22 surements with a low BRDF inversion quality (50% of more fill values) are excluded from the
23 analysis. The effect of excluding measurements with low BRDF inversion quality (e.g., 50% of
24 more fill values) on the results has been studied (results not shown) and did not influence the
25 effect of MAP and TAC (see Section 4). The 8-day BRDF corrected NDVI values are aggregated
26 to monthly mean NDVI values to facilitate comparison with other data sets and reduce noise due
27 to BRDF and cloud effects^{2,3}.

28 The MODIS NDVI data, despite the corrections and temporal compositing, still contains
29 residual invalid measurements (e.g., cloud effects). The quality flags are used to remove all
30 invalid observations and also derive the percentage of invalid measurements per pixel through
31 the whole time span as a summary of cloud and atmospheric effects. We include the maximum
32 number of consecutive and percentage of invalid measurements in our modeling approach to

33 account for the influence of clouds and atmospheric effects (Section 3).

34 Vegetation Optical Depth data (VOD) from the Advanced Microwave Scanning Radiome-
35 ter - Earth Observing System (AMSR-E, July 2002–2011, 0.25 degree spatial resolution) are
36 used⁴. VOD represents water content in aboveground woody and leaf biomass and is sensitive
37 to long-term climate changes. Unlike NDVI, the microwave VOD are least affected by atmo-
38 spheric and weather conditions^{4,5}. Also, satellite microwave observations have shown that forest
39 in western Amazonia experienced a strong water deficit during the dry season of 2005 and the
40 slow recovery (> 4 year) of forest canopy structure⁶. The presence of open water affects the
41 microwave emissions and may lead to underestimates of VOD values⁷. Because of this, regions
42 with extensive lakes, reservoirs, rivers and flooded vegetation were masked out^{4,8}.

43 In summary, the NDVI, operating in the optical regime, is sensitive to chlorophyll abun-
44 dance and photosynthetically active biomass of the leaves, whereas the microwave-based VOD is
45 an indicator of the vegetation water content in total above ground biomass, i.e., including wood
46 and leaf components⁴. Previous studies indicate that the fluctuations in VOD typically correlated
47 to precipitation variations, and that the mutually independent VOD and NDVI do not necessar-
48 ily respond in identical manners⁹. Considering both products together provides a more robust
49 assessment of long-term vegetation dynamics at the global scale.

50 1.2 Precipitation and temperature data

51 Mean annual precipitation (MAP) and mean annual temperature (MAT) from 1962–2011 is de-
52 rived from the *Climatic Research Unit* (CRU) monthly rainfall and minimum temperature data
53 set (<http://badc.nerc.ac.uk/browse/badc/cru>). Decelerating growth rates in tropical forest trees
54 have been reported to be linked with increasing annual mean daily minimum temperatures and
55 decreasing mean annual precipitation^{10,11}. The latest and currently longest global precipitation
56 product is used, i.e., CRU TS 3.20¹² (0.5° by 0.5°).

57 We also used precipitation data from the Tropical Rainfall Measuring Mission (TRMM;
58 3B42-v6; monthly, 0.25°) for the period of 1998–2011 to derive MAP at a higher spatial resolution
59 when compared to the CRU data set. The satellite observations of rainfall from TRMM has been
60 able to capture patterns of rainfall magnitude and seasonality undetected by CRU due to lack of
61 operating ground stations, particularly over the past decade^{13,14} (Figure S1).

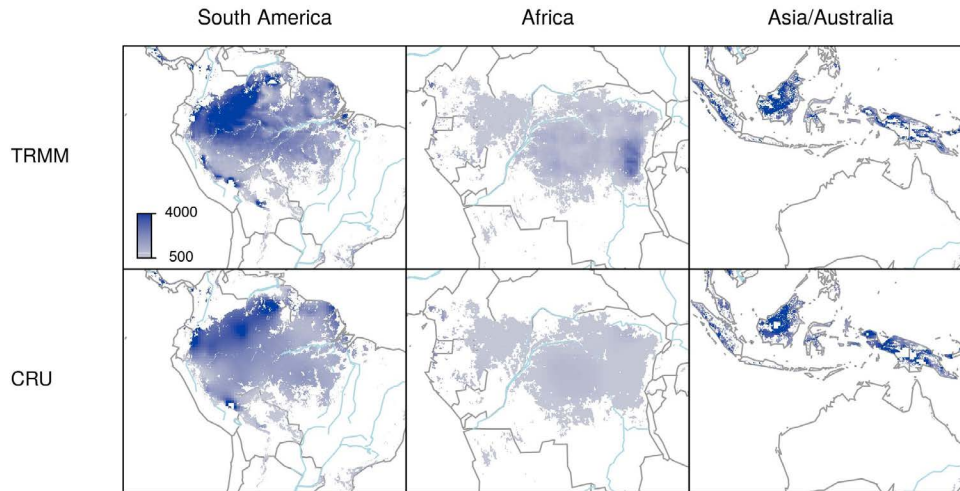


Figure S1: Mean Annual Precipitation (MAP, mm/year) derived from monthly TRMM (2000–2011) and CRU (1962–2011) data sets.

62 1.3 Soil fertility data

63 In addition, as a potential explanation of spatial patterns of slowing down of vegetation activ-
 64 ity in the pan-tropics, the Total Exchangeable Bases (TEB) of the top soil, as provided by the
 65 harmonised world soil database¹⁵, is included in the model (see Figure S3 and Section 3).

66 1.4 Sampling strategy for selecting intact and undisturbed evergreen trop- 67 ical forests

68 The centroid of each MODIS raster cell is used to sample NDVI time series, percentage tree
 69 cover and MAP, soil TEB for the intact and undisturbed evergreen forests in the tropics (35° S and
 70 15° N). The sampled data is then used to assess the continent specific relationship between NDVI
 71 based measures of critical slowing down and environmental variables (e.g., MAP). Figure S1
 72 illustrates the mean annual precipitation derived from monthly TRMM (2000–2011) and CRU
 73 (1962–2011) data in the pan-tropics.

74 We concentrated the analysis on forested areas with tree cover higher than 60%¹⁶ based on
 75 the MODIS percentage tree cover product and selected intact and undisturbed evergreen trop-
 76 ical forests of Africa, South America and South East Asia. We selected intact forest using
 77 the intact forests reported by the World Intact Forest Landscape (IFL) map ([http://www.
 78 intactforests.org/](http://www.intactforests.org/),¹⁷).

79 We selected only evergreen tropical forest and excluded the human impacted, bare or
80 flooded areas reported by the Global Land Cover 2000 product (GLC2000,¹⁸). The percent-
81 age tree cover for the year 2010 is derived from MODIS vegetation continuous fields product
82 (MOD44B, version 5, spatial resolution 250m). The world's IFL map is a spatial database (scale
83 1:1,000,000) that shows the extent of the intact forest landscapes (IFL) for year 2000. Intact
84 forests are selected when more than 90% of the MODIS pixels are intact as indicated by the IFL
85 raster layer. Undisturbed evergreen forests are identified using the GLC2000 when less than 10%
86 of the pixels had experienced intensive human impacts (codes: 16–18, and 22) or were bare or
87 flooded (codes: 15 and 19–21)¹⁹.

88 2 Indicators of slowing down

89 We employ the temporal autocorrelation (TAC) of the NDVI time series as the main indicator of
90 slowing down. As TAC varies substantially over space and time, the TAC is determined individu-
91 ally for each pixel so that the resulting estimates cannot be confounded with spatial correlations.
92 Moreover, we account for the possibility that the TAC itself varies over time while adjusting long-
93 term trends and seasonal periodic patterns. Specifically, each pixel is first detrended (including
94 seasonal adjustment) and subsequently the TAC of the detrended NDVI is computed using two
95 different methods.

96 2.1 Detrending

97 In order to obtain an accurate estimate of TAC , time series need to be stationary without long-
98 term trends or seasonal periodic patterns^{20,21} (just called “trends” in the following). Otherwise, if
99 trends are not accounted for, the TAC estimate may be biased. Fortunately, deterministic trends
100 can be removed efficiently without the need for specifying the TAC pattern (see Fuller, 1996,
101 pp. 476–480)²².

102 To select a detrending method that is most appropriate for time series with potentially
103 time-varying seasonal amplitude (i.e., size of the seasonal effect) and phase (i.e., start of the
104 season) we conducted an extensive simulation study (paper in preparation). In the simulation
105 study, we assessed how well various methods from the literature can detrend artificial time se-
106 ries with typical characteristics of NDVI and VOD time series data^{23,24}. More precisely, we
107 considered artificial time series with all combinations of several long-term trend patterns (none,
108 linear, smoothly varying), seasonal patterns (none, cyclical, time-varying amplitude and phase),
109 and TAC patterns (none, constant, linear, smoothly varying). The considered detrending methods
110 are: linear filtering (Holt-Winters with additive or multiplicative season)^{25,26}, Kalman filtering

111 using state space models (with potentially time-varying trend and seasonal pattern)²⁷, linear re-
112 gression (linear trend with harmonic seasonal pattern), additive regression (smooth long-term
113 trend with either smooth cyclical or time-varying seasonal pattern)^{28–30}, and LOESS smoothing
114 (with smooth long-term trend and either stable or smoothly varying seasonal pattern)³¹.

115 The result of the simulation study illustrates that many of the various detrending methods
116 lead to qualitatively similar results and can reliably recover both the trend and the TAC. How-
117 ever, not surprisingly, if amplitude/phase are indeed time-varying in the simulation, methods that
118 allow for such flexibility perform clearly better. To make sure that our results are robust against
119 many conceivable trend patterns, we employ three particularly well-performing models in the
120 following: season-trend decomposition by LOESS smoothing (STL)³¹, decomposition based on
121 additive models including time-varying harmonic seasonal effects (STA)³⁰ and extended Kalman
122 filtering of state space model (SSM)²⁷ with time-varying parameters to capture smoothly varying
123 and dynamically changing trends, respectively. An advantage of the STA and SSM when com-
124 pared to the STL approach is that no interpolation of missing data is required. Furthermore, to
125 ensure stationarity of the detrended series KPSS tests³² are carried out.

126 Figure S2 illustrates the ability of the detrending methods used to model highly complex
127 time trends and seasonal patterns. Figure S2 illustrates that all, slowly and smoothly vary periodic
128 patterns have been removed prior to computing TAC. What the Figure S2 shows is that here dryer
129 regions tend to have higher TAC (i.e. slower recovery times after disturbance). The key point
130 here is that the TAC has been estimated from time series from which all potential deterministic
131 patterns, i.e. long-term trends and time-varying seasonality have been removed. All potentially
132 occurring trend and seasonal variation has been removed using an “ensemble” of detrending
133 methods to enable a “true” TAC estimate^{21,33}.

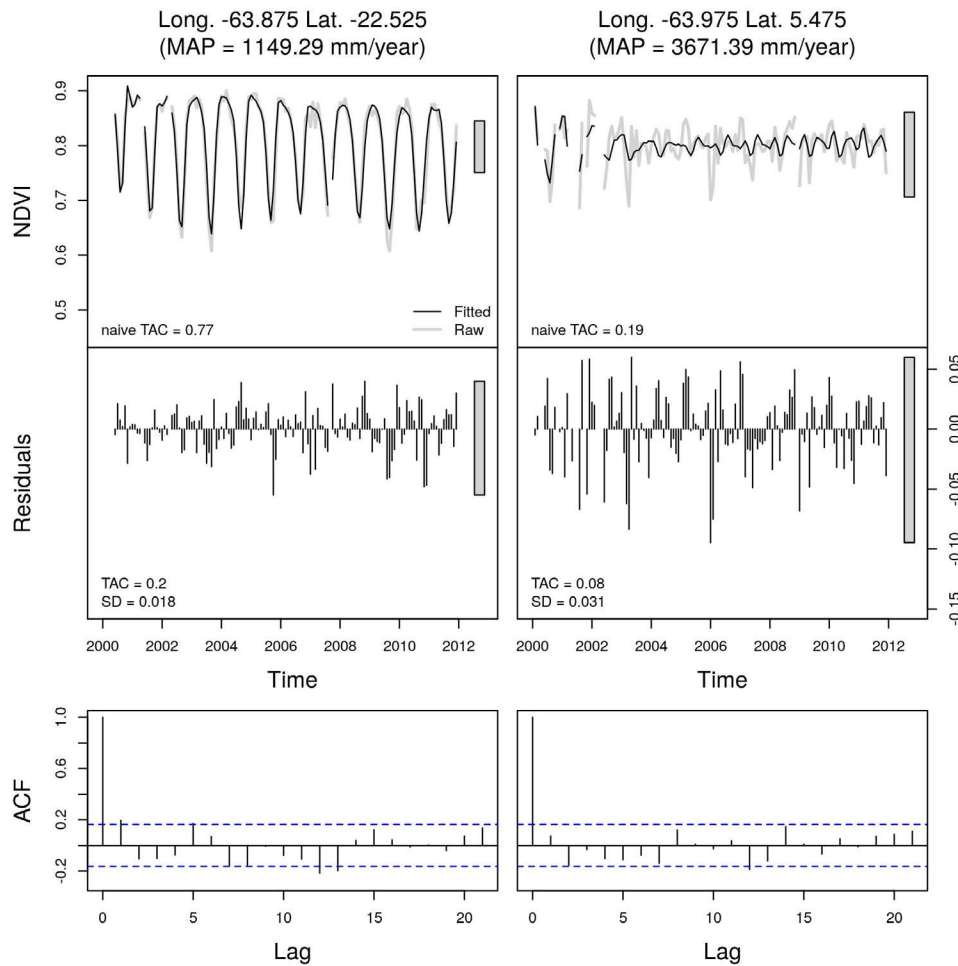


Figure S2: Examples of detrending of undisturbed forest NDVI time series using the STA in South America. The left column represents a time series of a pixel with low MAP and pronounced seasonality, the right column high MAP and less pronounced seasonality. The top row shows the raw NDVI time series in gray versus the fitted times series from the STA in black. Both cases suggest accurate detrending although both time series exhibit time-varying seasonal patterns and data gaps. The residuals indicate stationarity which is also confirmed by the KPSS test. The estimated ACF1 at higher lags also suggests no remaining significant seasonal patterns. The reported TAC estimates derived from the ACF1 method are higher for the drier forest pixel. The SD results are rather similar. The corresponding interquartile range of the estimated 2 year ACF1 rolling method is 0.15–0.31 (mean 0.18) for the low MAP pixel and 0.09–0.15 (mean 0.09) for high MAP pixel.

134 2.2 TAC estimation

135 After removing long-term trends and seasonal patterns individually for each pixel, TAC and SD
 136 can simply be estimated by the corresponding full-sample estimators. In case of TAC, we simply
 137 employ the empirical autocorrelation function at lag 1 (ACF1). Moreover, to account for possible
 138 time-varying effects in TAC that could affect TAC over the whole observation period, we applied
 139 a two year rolling window ACF1 approach²⁰.

140 Combining the three detrending methods (STL, STA and SSM) with the two TAC esti-
 141 mators above yields six possible combinations (see Table S1). In all subsequent analyzes we
 142 always report results of this ensemble of six method combinations in order to demonstrate that
 143 all conclusions are robust against the methods for detrending and TAC estimation.

Detrend method	TAC method
STL	ACF1
STL	ACF1 roll
STA	ACF1
STA	ACF1 roll
SSM	ACF1
SSM	ACF1 roll

Table S1: Overview of all combinations of the detrending and TAC methods.

144 3 Models

145 To avoid potential bias in TAC and SD estimates (Section 2) when the time series contain a
 146 large number of missing values by the exclusion of bad quality measurements (Section 1.1),
 147 strict quality criteria are used to exclude time series with a percentage of missing values (further
 148 referred to as %NAs) larger than 10%, as well as time series containing more than 2 consecutive
 149 missing values, i.e., maximum of 2 months, further referred to as NA_{maxgap} . To further reduce
 150 the influence by clouds and haze, locations with mean annual precipitation (MAP) of more than
 151 4000 mm are excluded similar to Hirota et al.¹⁶. The total number of time series analysed for the
 152 different satellite data sets and detrending methods is shown in Table S2.

153 We assessed the relationship of TAC and SD of intact evergreen forest pixels with environ-
 154 mental covariates using additive regression models^{28,34,35}. Since TAC and SD estimates may vary
 155 by the detrending and TAC methods used, we estimated an ensemble of the 6 different detrending
 156 and TAC method combinations (see Table S1) for each of the two data sets (see Table S2).

Data set	Continent	# Obs. STL	# Obs. STA	# Obs. SSM
MODIS NDVI	South America	41492	40833	40806
	Africa	21701	21511	21501
	Asia/Australia	7630	7324	7318
VOD	South America	4897	4897	4897
	Africa	889	889	889
	Asia/Australia	374	374	374

Table S2: Number of time series observations (# Obs.) analysed for different continents, data sets (MODIS NDVI 2000–2011, VOD 2002–2011), and detrending methods (STL, STA and SSM).

157 The main explanatory variable of interest is MAP derived from TRMM data (see Sec-
 158 tion 1.2). To assess the influence of other environmental variables like percentage tree cover
 159 (%Trees), the effect of varying seasonality (Season, estimated as the amplitude of the seasonal
 160 component using STL decomposition³¹), TAC and SD of precipitation time series (PreTAC,
 161 PreSD, derived from the detrended precipitation time series), mean annual temperature (MAT),
 162 the soil fertility (Soil) or by %NAs (as a proxy for cloud cover), we estimated additive models
 163 given by

$$TAC = \gamma_0 + f_1(MAP) + f_2(\%Trees) + f_3(Season) + f_4(PreTAC) + f_5(MAT) + f_6(Soil) + f_7(\%NAs) + f_8(Long, Lat) + \epsilon \quad \epsilon \sim N(0, \sigma^2), \quad (1)$$

164 where for models with SD as the response variable PreTAC is exchanged by PreSD. Also
 165 note that the effect of %NAs is not included using VOD data, since the time series do not contain
 166 missing values. The parameter γ_0 is the usual model intercept, the functions f_1, \dots, f_7 are smooth
 167 functions and are modeled by splines. Furthermore, function f_8 accounts for spatial effects of
 168 longitude and latitude pixel coordinates on TAC or SD, i.e., the function models spatial variation
 169 due to factors like dry season length, fire frequency, topography and potential herbivory effects,
 170 that have not been measured¹⁹.

171 4 Results

172 When computing the KPSS test on a 5% significance level, all detrending methods generated
 173 stationary time-series. Table S3 shows the percentage rates of pixels that according to the KPSS
 174 test are stationary. 98% of all VOD and 93-98% of all MODIS time-series selected for analysis
 175 (see sampling strategy) are stationary.

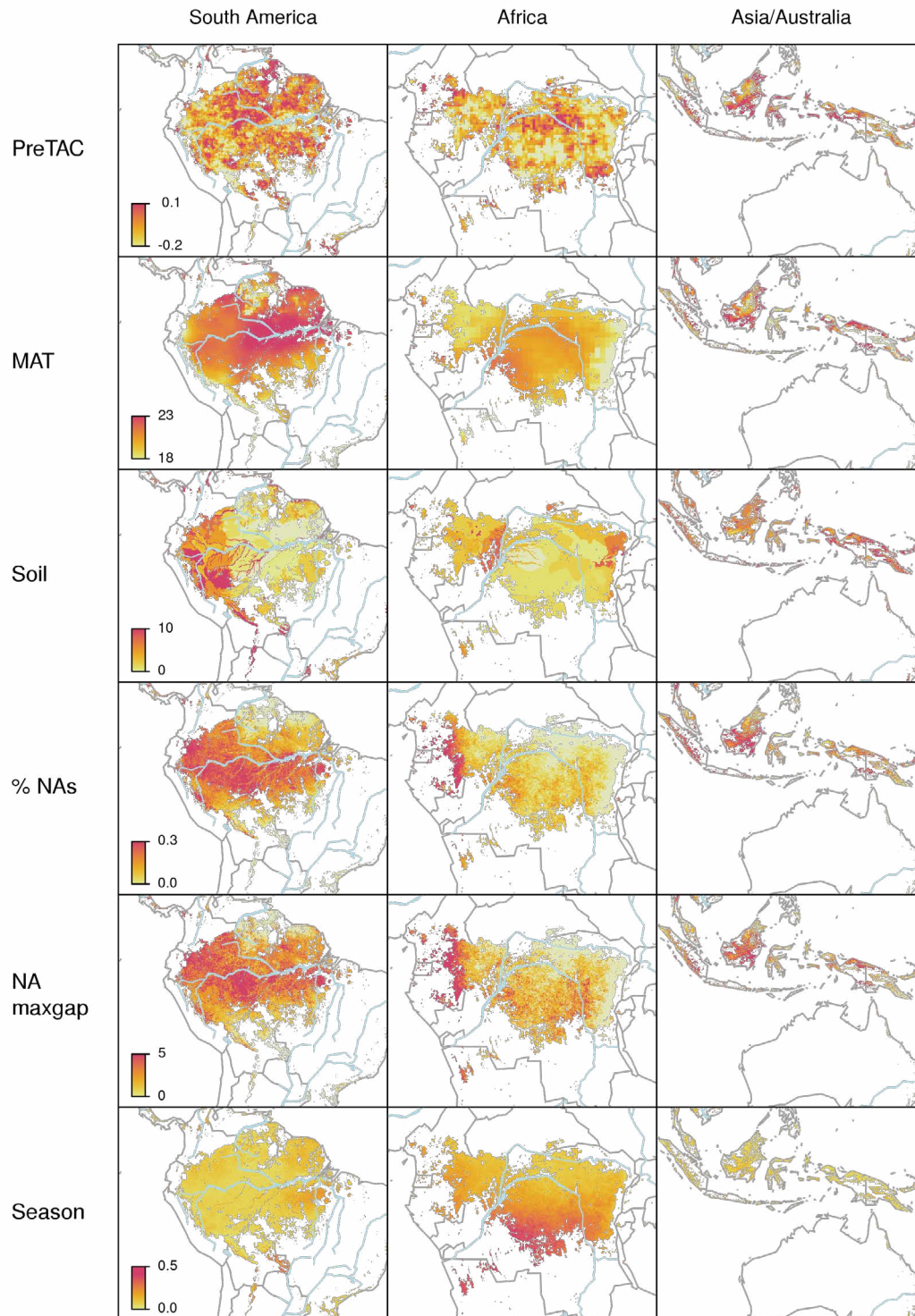


Figure S3: Continent-specific maps of detrended TRMM precipitation temporal autocorrelation (PreTAC), mean annual CRU temperature (MAT), soil fertility (Soil), percentage missing observations (%NAs), maximum number of consecutive missing observations (NA maxgap) in the MODIS NDVI time series and seasonal amplitude of the NDVI time series (Season).

Detrend method	KPSS test
VOD SSM	0.989
VOD STL	0.990
VOD STA	0.986
MODIS SSM	0.932
MODIS STL	0.981
MODIS STA	0.935

Table S3: Percentage rates of pixels that according to the KPSS test are stationary at a 5% significance level. The percentages are calculated for selected MODIS and VOD based time series following the sampling strategy described above.

176 4.1 Environmental effects on TAC

177 Models are fitted per continent for all combinations of detrending and TAC methods (see Sec-
178 tion 2, Table S1 and Section 3) to study the influence of MAP on TAC.

179 The MAP effect on TAC is similar across continents and satellite derived MODIS NDVI and
180 VOD data sets (Figures S4 and S5). It is shown that for a decreasing MAP the TAC increases while
181 accounting for a all other relevant environmental variables (%Trees, Soil, MAT), and potential
182 data driven effects like seasonality (Season) and data quality (%NAs) of the NDVI signal and
183 TAC derived from precipitation (PreTAC). The MAP and Season are the two most important
184 variables and show that for decreasing MAP (i.e., increasing dryness) the TAC increases while
185 seasonality increases (e.g., forest types showing more seasonality). This confirms that dryer
186 evergreen forest are slowing down while showing more seasonal variability for decreasing tree
187 cover percentages (%Trees).

188 Across different continents, the temporal autocorrelation of precipitation (PreTAC) has
189 no effect on TAC. Furthermore, it is shown that TAC is increasing with increasing MAT (the only
190 exception being the NDVI pattern for Africa, Figure S4). This could illustrate earlier reported
191 findings that decelerating growth rates in tropical forest trees are linked with increasing mean
192 annual temperature^{10,11}.

193 The estimated ensemble mean levels of TAC (the intercept γ_0 in Eq. 1) for all data sets and
194 continents are shown in Table S4. According to MODIS and VOD data, the estimates suggest
195 that the mean level of TAC is the lowest in Africa.

Data set	Continent		
	South America	Africa	Asia/Australia
MODIS NDVI	0.15 (0.00046)	0.08 (0.00057)	0.18 (0.00109)
VOD	0.19 (0.00151)	0.13 (0.00298)	0.14 (0.00512)

Table S4: Estimated ensemble mean (intercepts γ_0 in Eq. 1) and standard error of the derived TAC of all data sets, detrending and TAC methods.

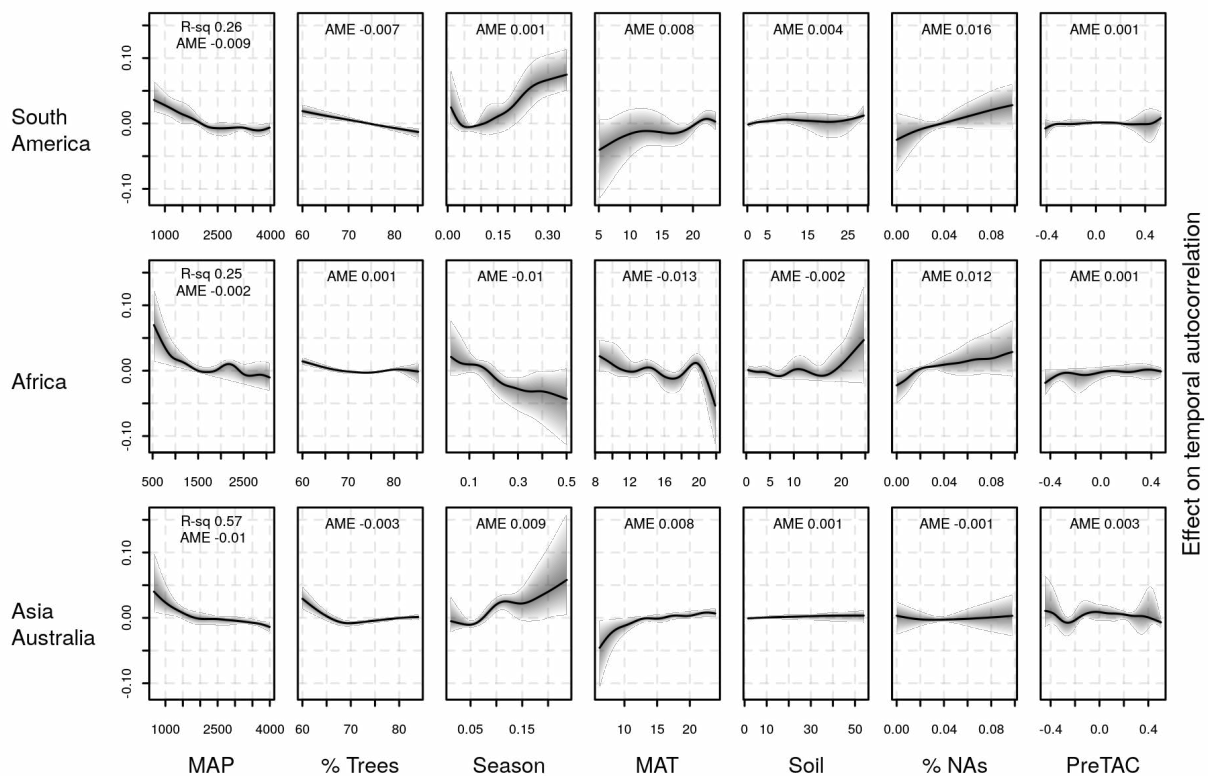


Figure S4: Estimated effects for all terms of model (Eq. 1) on TAC using NDVI MODIS (2000–2011), derived from different detrending and TAC methods. The black solid lines represent the mean curve from an ensemble of six curves (i.e., three detrending methods and two indicators of temporal autocorrelation). The gray shaded area highlights the range of the six individual ensemble members. The average R^2 and Average Marginal Effect (AME) are reported in the top range of the plots.

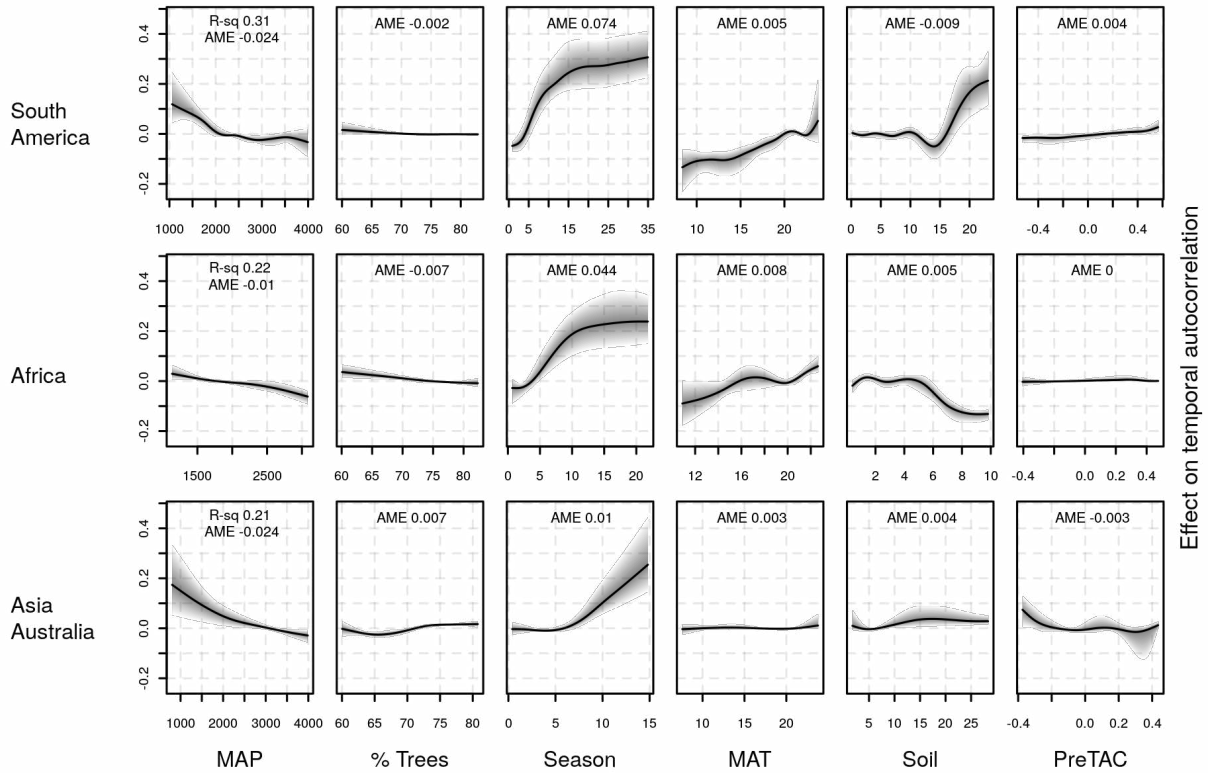


Figure S5: Estimated effects for all terms of model (Eq. 1) using the VOD (2002-2011), derived from different detrending and TAC methods. The black solid lines represent the mean curve from an ensemble of six curves (i.e., three detrending methods and two indicators of temporal autocorrelation). The gray shaded area highlights the range of the six individual ensemble members. The average R^2 and Average Marginal Effect (AME) are reported in the top range of the plots.

196 **4.2 Environmental effects on SD**

197 Figures S6 and S7 illustrate that the MAP and other environmental variables except the varying
 198 seasonality (Season), have a minor effect on SD variation as shown in the y-axis. The mod-
 199 els explaining SD variability do not show clear patterns for the MAP effect as we obtain when
 200 modeling TAC.

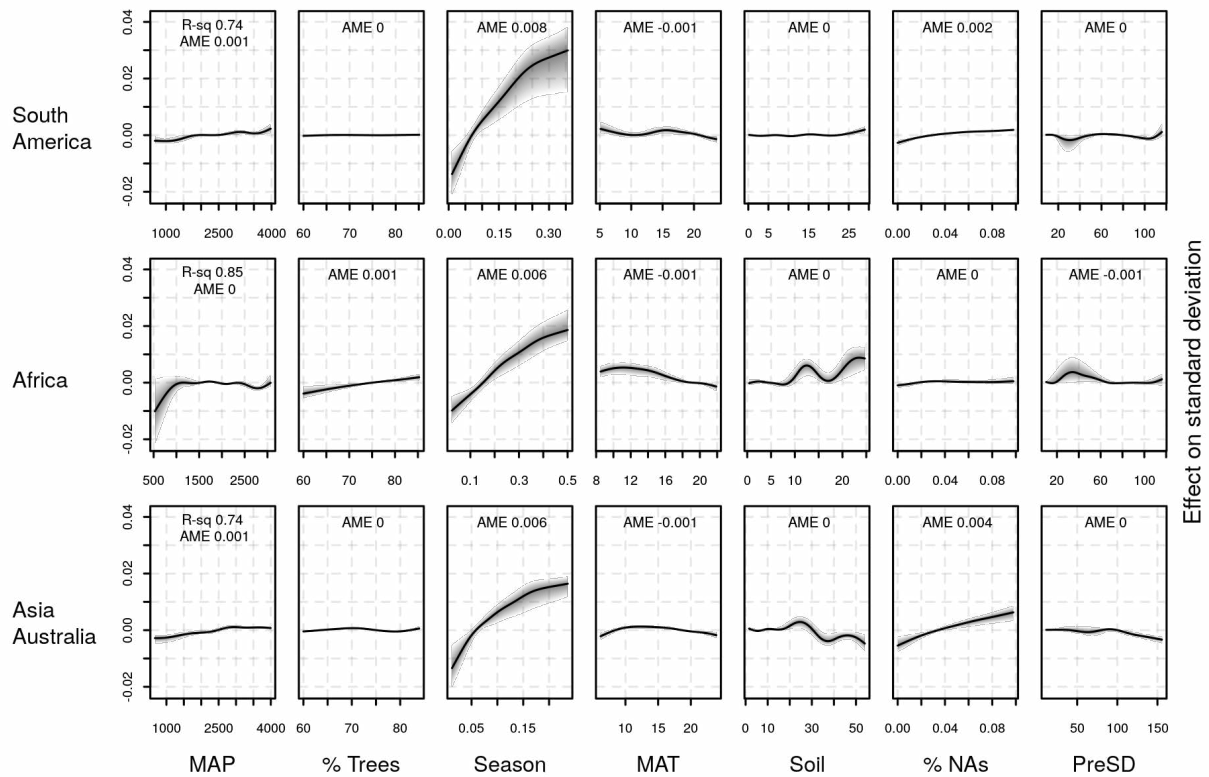


Figure S6: Estimated effects for all terms of model (1) on SD using NDVI MODIS (2000–2011), derived from different detrending methods. The black solid lines represent the mean curve from an ensemble of six curves (i.e., three detrending methods and two indicators of temporal auto-correlation). The gray shaded area highlights the range of the six individual ensemble members. The average R^2 and Average Marginal Effect (AME) are reported in the top range of the plots.

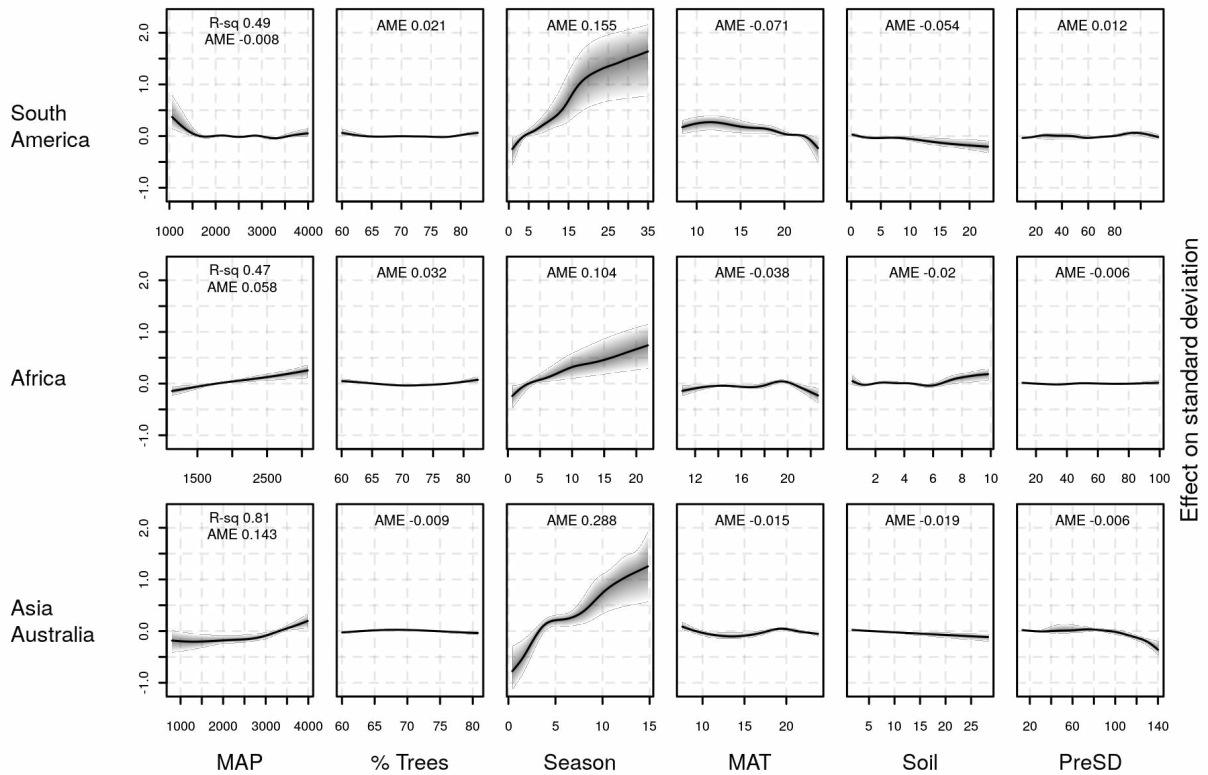


Figure S7: Estimated effects for all terms of model (1) using VOD (2002-2011), derived from different detrending methods. The black solid lines represent the mean curve from an ensemble of six curves (i.e., three detrending methods and two indicators of temporal autocorrelation). The gray shaded area highlights the range of the six individual ensemble members. The average R^2 and Average Marginal Effect (AME) are reported in the top range of the plots.

201 **5 Software**

202 The analysis is fully processed within the statistical environment R³⁶ version 3.1.0. For satellite
203 image manipulation we used the **raster**³⁷ and **bfast**²³ package ([http://bfast.R-Forge.
204 R-project.org/](http://bfast.R-Forge.R-project.org/)). Detrending using the state-space model is based on the **sspir**^{38,39} package,
205 all models were fitted by the **mgcv**⁴⁰ package.

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210 data set (2002-2011).

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