## **Supplementary information:** Low efficiency of large volcanic eruptions in transporting very fine ash into the atmosphere

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Very fine ash partitioning  $(\epsilon = Q_a/Q_s)$ 

**Supplementary Information Figure S1. Summary of the main findings.** We show that the partitioning ( $\varepsilon$ ) between the amount of tephra injected into the volcanic plume and the fine ash fraction that survive proximal sedimentation forming distal ash clouds is highly variable. During sustained eruptions (Plinian, Subplinian and Small/Moderate styles) this partitioning inversely scales with the eruption intensity ( $Q_s$ ). Actually, the more the plume is concentrated in fine ash, the lower the proportion of fine ash injected into the atmosphere. This is due to the efficiency of collective particle settling in fine ash-rich plumes (such as Plinian and Subplinian) which enhance early and en masse fallout of fine ash, hence leading to very low  $\varepsilon$  values. We thus design a style-related operational parameterization using our partitioning coefficients ( $\varepsilon$ =0.5% for Plinian,  $\varepsilon$ =0.8% for Subplinian, and  $\varepsilon$ =3.2% for Small/Moderate) to be implemented as input parameters for ash-cloud-dispersal models run by the 9 VAACs (Volcanic Ash Advisory Centres) worldwide for the production of ash forecast maps.



**Supplementary Information Figure S2. Atmospheric ash concentration simulations.** The four simulations are produced by the volcanic ash dispersion model MOCAGE of the Toulouse VAAC based on two eruptive scenarios, using different partitioning coefficients and present-day meteorological data. **a**, Simulation of ash dispersion in the atmosphere at Etna volcano 7 hours after a Small/Moderate eruption based on the 27 October, 2002 event (Supplementary Information Table 2), using the VAAC-default  $\varepsilon$  value of 5%. **b**, Same simulation conditions and scenario but using the Small/Moderate  $\varepsilon$  value established in this study at 3.2%. The extent of the No-Fly zone (4 mg/m<sup>3</sup> for an ash cloud 500-m thick<sup>\*</sup>) is overestimated by the VAAC-default  $\varepsilon$  value. **c**, Simulation of ash dispersion in the atmosphere at Cordón Caulle volcano 14 hours after a Subplinian eruption based on the 4 June, 2011 event (Supplementary Information Table 2), using the VAAC-default  $\varepsilon$  value of 5%. **d**, Same simulation conditions and scenario but using the Subplinian  $\varepsilon$  value established in this study at 0.8%. The extent of the No-Fly zone (4 mg/m<sup>3</sup> for an ash cloud 500-m thick<sup>\*</sup>) is overestimated by the value of 5%. **d**, Same simulation conditions and scenario but using the Subplinian  $\varepsilon$  value established in this study at 0.8%. The extent of the No-Fly zone (4 mg/m<sup>3</sup> for an ash cloud 500-m thick<sup>\*</sup>) is greatly overestimated by the VAAC-default  $\varepsilon$  value. <sup>\*</sup>The threshold at 4 mg/m<sup>3</sup> was first established by European commission after the Eyjafjallajökull 2010 eruption. It is now described by EASA (European Aviation Safety Agency) and used in the emergency plan EUR/NAT (EURopean and North ATlantic office) as the "High" contamination level.



Supplementary Information Figure S3. Evaluation of the quality of the tephra deposition simulations of the  $23^{rd}$  February 2013 Etna eruption. The loading of tephra on the ground (x-axis) was measured at individual locations in the field, while the computed loading of tephra deposited on the ground (y-axis) at the same locations was generated by the FALL3D tephra-deposition model. All values and locations are listed in Supplementary Information Table 3. The filled circles correspond to the results of simulation 1 (Fig. 4a of the main text) using an input  $Q_s$  estimated with our satellite-derived statistical model (see equation in legend for low SiO<sub>2</sub> content and open system). In this case, all values fall within an error envelope of a factor of 5, except at location #7. At locations #1, 4, 5, 9 and 10 there is a perfect agreement with measured loading. The open circles correspond to the results of simulation 2 (Figure 4b of the main text) using an input  $Q_s$  estimated with the standard empirical scaling law (see equation in legend). Simulation 2 shows a larger discrepancy with the measured loading at location #1, 4, 5, 9 and 10, with an error factor larger than one order of magnitude.

	Kelut	Cordón Caulle	Etna
Eruptive style	Plinian	Subplinian	Small/Moderate
Eruption duration (hours)	3	24	10
Plume height (km)	20	13.7	3.65
Qs (kg/s)	6.02E+07	7.00E+06	2.42E+04
VAAC partitioning (ɛ)	5	5	5
Eruptive-style partitioning ( $\epsilon$ )	0.45	0.8	2.23
TGSD (μm)	0.1-100*	0.1-100*	0.1-100*
Column mass distribution	Uniform	Uniform	Uniform
Meteo data resolution	0.5°x0.5°	0.5°x0.5°	0.5°x0.5°
Wind field time (T <sub>0</sub> )	20170620 1800 UTC	20170620 1900 UTC	20170621 0100 UTC

Supplementary Information Table S1. Parameters of 3 eruptive scenarios for MOCAGE simulations (Toulouse VAAC). The 3 scenarios cover each eruptive style for sustained eruptions (Plinian, Subplinian and Small/Moderate) with related parameters used at the input of the MOCAGE model run by Toulouse VAAC. For each scenario, two simulations are computed using alternatively the VAAC operational partitioning coefficient (5%) and our modified partitioning coefficient depending of eruptive styles. \*The Total Grain Size Distribution (TGSD) ranges from 0.1-100  $\mu$ m and is split into 6 classes with 70% of the mass fraction within the 10-30  $\mu$ m size range<sup>48</sup>.

ld	Sampling Location	Distance (km)	Measured Load (kg/m2)	Measured Grain Size distribution mode (Φ)	Computed load Simulation 1	Computed load Simulation 2	
1	Baracca	5.35	2.10E+01	-3.5	2.20E+01	1.30E+00	
2	Casetta	5.45	5.90E+00	-4	2.20E+01	1.30E+00	
3	Bivio-007	5.79	5.50E+00	-4	2.30E+01	1.40E+00	
4	Forestale	7.32	2.20E+01	-3.5	2.50E+01	1.50E+00	
5	Chalet	10.21	3.20E+01	-2.5	2.60E+01	1.60E+00	
6	Castiglione	15.52	5.20E+00	-1.5	2.10E+01	1.30E+00	
7	Linguaglossa	15.6	1.20E+00	-3	2.00E+01	1.20E+00	
8	Messina	69.55	2.90E-01	1	1.30E+00	7.90E-02	
9	Cardinale	157.17	1.30E-02	2	1.50E-02	9.00E-04	
10	Brindisi	408.91	1.40E-03	3	2.00E-03	1.20E-04	

Supplementary Information Table S2. Field data and results of the tephra deposition simulations of the 23 February Etna eruption. Simulations run with the volcanic tephra-deposition model FALL3D, using an input  $Q_s$  estimated with our satellite-derived statistical model (simulation 1, See Figure 4a in the main text and Supplementary Information Fig. 4), and an input  $Q_s$  estimated with the standard empirical scaling law (simulation 2, See Figure 4b in the main text and Supplementary Information Fig. 4).

Qs (kg/s)		H (km a.v.)	TTest	Corr.	Bias	RMSE <sub>1</sub>	RMSE <sub>2</sub>	RMSE₃	К	k	TEM (kg)
Satellite-derived stat. model	2.75×10 <sup>6</sup>	5.5	0.6	1.0	0.1	0.37	5.33	1.77	0.41	2.48	1.09×10 <sup>10</sup>
Empirical scaling law	1.66×10 <sup>5</sup>	5.5	0.2	1.0	-0.2	0.45	0.81	0.73	6.85	2.48	6.58×10 <sup>8</sup>

Supplementary Information Table S3. Statistical metrics evaluating the goodness-of-the-fit of the tephra deposition simulations of the 23<sup>rd</sup> February Etna eruption. *Corr* is the correlation coefficient between the observed and the simulated tephra loading values (Obs and Sim, respectively), expressed as  $Corr(Obs, Sim) = \frac{Cov(Obs,Sim)}{\sigma_{Obs}\sigma_{Sim}}$ , where, Cov is the covariance coefficient and  $\sigma$  is the standard deviation. *Bias* is expressed as  $Bias = \frac{\sum_i Sim_i - Obs_i}{\sum_i Obs_i}$ , where i represents the i<sup>th</sup> individual tephra sample. *TTest* is calculated as  $TTest = \frac{\overline{Obs} - \overline{Sim}}{\sqrt{\frac{\sigma_{Obs}^2 + \sigma_{Sim}^2}{N}}}$ , where  $\overline{Obs}$  and  $\overline{Sim}$  are the means of the observed and simulated tephra loadings, respectively, and N is the number of samples. The TTest is a probabilistic coefficient that indicates how along the simulated and observed tephra loading uplues are. The three PMSEs are calculated using different

simulated tephra loadings, respectively, and N is the number of samples. The TTest is a probabilistic coefficient that indicates how close the simulated and observed tephra loading values are. The three RMSEs are calculated using different hypotheses regarding the uncertainty on the ash loading values.  $RMSE_1$  is calculated assuming that all the ash loading measurements have an equal uncertainty.  $RMSE_2$  and  $RMSE_3$  are calculated assuming a linear and uniform distribution, respectively, of the uncertainty on the ash loading measurements.

		Plinian	Subplinian	Small/Moderate
	phi [-log2(D in mm)]	wt.%	wt.%	wt.%
	-6	8.5E-01	2.2E-01	2.4E-04
	-5	1.8E+00	9.4E-01	1.7E-02
	-4	3.0E+00	2.9E+00	4.4E-01
	-3	4.4E+00	6.3E+00	4.4E+00
	-2	5.4E+00	9.8E+00	1.7E+01
	-1	5.8E+00	1.1E+01	2.6E+01
	0	6.1E+00	1.1E+01	1.7E+01
	1	8.3E+00	1.1E+01	9.9E+00
Total Grain Size	2	1.4E+01	1.4E+01	1.0E+01
Distribution (TGSD)**	3	1.9E+01	1.5E+01	8.7E+00
	4	1.7E+01	1.1E+01	4.8E+00
	5	1.0E+01	4.9E+00	1.7E+00
	6	3.7E+00	1.4E+00	3.7E-01
	7	8.6E-01	2.5E-01	5.0E-02
	8	1.3E-01	2.9E-02	4.2E-03
	9	1.2E-02	2.0E-03	2.2E-04
	10	7.3E-04	8.9E-05	7.5E-06
	11	3.6E-05	2.4E-06	1.6E-07
	12	2.1E-06	4.2E-08	2.0E-09
Total ash fraction	d<63µm (%)	31.6	17.4	6.9
Distal ash fraction	ε (%)	0.5	0.8	3.2
MER (Qs, kg/s)	closed-conduit	25,95Qa <sup>0,62*</sup> H <sup>1,95</sup>	25,95Qa <sup>0,62*</sup> H <sup>1,95</sup>	25,95Qa <sup>0,72*</sup> H <sup>1,95</sup>
	open-conduit	25,95Qa <sup>0,62*</sup> H <sup>1,4</sup>	25,95Qa <sup>0,62*</sup> H <sup>1,4</sup>	25,95Qa <sup>0,72*</sup> H <sup>1,4</sup>

Eruptive scenarios

\*0,62 is for High-SiO<sub>2</sub> content and 0,72 is for low-SiO<sub>2</sub> content of the magma \*\* TGSD calculated<sup>97</sup> using magma viscosity (Log<sub>10</sub> Pa.s) of 3,4 and 5 with plume height (km) of 5,15, and 25 for Small/Moderate, Subplinian, and Plinian eruptive scenarios, respectively.

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